

Evapotranspiration – Equations

Penman – Monteith (1965)

$$\lambda ET = \frac{\Delta(R_n - G) + \rho_a c_p \frac{(e_s - e_a)}{r_a}}{\Delta + \gamma(1 + \frac{r_s}{r_a})} = \frac{\Delta(Rn - G) + \rho_a c_p \frac{VPD}{r_a}}{\Delta + \gamma(1 + \frac{r_s}{r_a})} \quad (1)$$

Where:

λ is latent heat of vaporization of water (MJ kg⁻¹);

ET is evapotranspiration (mm d⁻¹);

R_n is net radiation at the crop surface (MJ m⁻² d⁻¹);

G is soil heat flux (MJ m⁻² d⁻¹);

ρ_a is air average density at constant pressure (kg m⁻³);

c_p is air specific heat (MJ kg⁻¹ °C⁻¹);

e_s is saturation vapor pressure (kPa);

e_a is actual vapor pressure (kPa);

$(e_s - e_a)$ is saturation vapor pressure deficit (kPa);

VPD is vapor pressure deficit (kPa);

Δ is slope vapor pressure curve (kPa °C⁻¹);

γ is psychrometric constant (kPa °C⁻¹);

r_s is surface resistance (s m⁻¹);

r_a is aerodynamic resistance (s m⁻¹).

 **Penman – Monteith (FAO 56)**

$$ET_o = \frac{0.408\Delta(R_n - G) + \gamma \frac{900}{T_{mean} + 273} u_2(e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (2)$$

Where:

ET_o is reference evapotranspiration (mm day^{-1});

R_n is net radiation at the crop surface ($\text{MJ m}^{-2} \text{ day}^{-1}$);

G is soil heat flux density ($\text{MJ m}^{-2} \text{ day}^{-1}$);

T is mean daily air temperature at 2 m height (m s^{-1});

u_2 is wind speed at 2 m height (m s^{-1});

e_s is saturation vapor pressure (kPa);

e_a is real vapor pressure (kPa);

$e_s - e_a$ is saturation vapor pressure deficit (kPa);

Δ is slope vapor pressure curve ($\text{kPa } ^\circ\text{C}^{-1}$);

γ is psychrometric constant ($\text{kPa } ^\circ\text{C}^{-1}$).

Considering the following restructure from Penman-Monteith (1965) equation:

$$\lambda ET = \frac{\Delta(R_n - G) + \rho_a c_p \frac{(e_s - e_a)}{r_a}}{\Delta + \gamma(1 + \frac{r_s}{r_a})} \quad (3)$$

$$\lambda ET \left(\Delta + \gamma(1 + \frac{r_s}{r_a}) \right) = \Delta(R_n - G) + \rho_a c_p \frac{(e_s - e_a)}{r_a} \quad (4)$$

$$ET \left(\Delta + \gamma(1 + \frac{r_s}{r_a}) \right) = \frac{\Delta(R_n - G) + \rho_a c_p \frac{(e_s - e_a)}{r_a}}{\lambda} \quad (5)$$

$$ET \left(\Delta + \gamma(1 + \frac{r_s}{r_a}) \right) = \frac{\Delta(R_n - G)}{\lambda} + \frac{\rho_a c_p \frac{(e_s - e_a)}{r_a}}{\lambda} \quad (6)$$

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + \frac{\rho_a c_p \frac{(e_s - e_a)}{r_a}}{\lambda}}{\Delta + \gamma(1 + \frac{r_s}{r_a})} \quad (7)$$

Considering the following approach of all parameters from Penman-Monteith (FAO56) equation:

"A hypothetical reference crop with an assumed crop height of 0.12 m, a fixed surface resistance of 70 s m⁻¹ and an albedo of 0.23." (Allen et al. 1998).

$$r_a = \frac{208}{u_2} \quad (8)$$

$$r_s = 70 \quad (9)$$

Replace the parameters from reference crop in PM equation:

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + \frac{\rho_a c_p \frac{(e_s - e_a)}{r_a}}{\Delta + \gamma(1 + \frac{r_s}{r_a})}}{(10)}$$

Considering:

$$\rho_a = \frac{P}{T_{kv} R_d} \quad (11)$$

$$c_p = \frac{\gamma \varepsilon \lambda}{P} \quad (12)$$

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + \frac{\frac{P}{T_{kv} R_d} \frac{\gamma \varepsilon \lambda}{P} \frac{(e_s - e_a)}{\frac{208}{u_2}}}{\Delta + \gamma(1 + \frac{70}{208})}}{u_2} \quad (13)$$

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + \frac{\frac{\gamma \varepsilon \lambda}{T_{kv} R_d} \frac{u_2}{208} (e_s - e_a)}{\Delta + \gamma(1 + 0.34 u_2)}}{\lambda} \quad (14)$$

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + \frac{\gamma \varepsilon}{T_{kv} R_d} \frac{u_2}{208} (e_s - e_a)}{\Delta + \gamma(1 + 0.34 u_2)} \quad (15)$$

Considering:

ε is ratio molecular weight of water vapor/dry air = 0.622.

T_{kv} is virtual temperature = $1.01(T + 273) \sim (T + 273)$.

R_d is universal gas constant specific for dry air = $0.287 \text{ kJ kg}^{-1} \text{ K}^{-1}$.

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + \frac{\gamma 0.622}{(T + 273)0.287} \frac{u_2}{208} (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (16)$$

Considering a time of one day (86400 seconds):

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + 86400 \frac{\gamma 0.622}{(T + 273)0.287} \frac{u_2}{208} (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (17)$$

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + \gamma \frac{900}{(T + 273)} u_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (18)$$

Considering the latent heat vaporization (λ) equal = 2.45 MJ kg⁻¹

“As λ varies only slightly over normal temperature ranges a single value of 2.45 MJ kg⁻¹ is taken in the simplification of the FAO Penman-Monteith equation. This is the latent heat flux for an air temperature of about 20 °C” (Allen et al. 1998).

$$ET = \frac{0.408\Delta(R_n - G) + \gamma \frac{900}{(T + 273)} u_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (19)$$

Considering the time of 1 hour (3600 seconds):

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + 3600 \frac{\gamma 0.622}{(T + 273)0.287} \frac{u_2}{208} (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (20)$$

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + \gamma \frac{37.51}{(T + 273)} u_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (21)$$

Considering the time of 30 minutes (1800 seconds):

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + 1800 \frac{\gamma 0.622}{(T + 273)0.287} \frac{u_2}{208} (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (22)$$

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + \gamma \frac{18.75}{(T + 273)} u_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (23)$$

Considering the time of 15 minutes (900 seconds):

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + 900 \frac{\gamma 0.622}{(T + 273)0.287} \frac{u_2}{208} (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (24)$$

$$ET = \frac{\frac{\Delta(R_n - G)}{\lambda} + \gamma \frac{9.37}{(T + 273)} u_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (25)$$

Reference:

Allen RG, Pereira LS, Raes D, Smith M (1998). Crop evapotranspiration - Guidelines for computing crop water requirements - FAO Irrigation and drainage paper 56. FAO, Rome, 300(9), D05109.

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